# Geophysical investigation of the Groundwater Potentials in the Imo Shale Formation of Ehime Mbano Area, Southern Nigeria

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#### **ABSTRACT**

This study was carried out to investigate the groundwater potential in some selected villages in the Imo Shale Formation of Ehime Mbano Area, Southern Nigeria. A total of ten (10) vertical electrical soundings were conducted using the ABEM SAS 4000 Terrameter. The data obtained were subjected to interpretation by partial curve matching and then by computer iteration. A total of ten geoelectric layers were delineated in this study. Results show that the aguifer is located within the fifth layers at an average depth of about 31 m - 40 m for VES locations 1 - 4; while the resistivity of the aguifer ranged between 1240 Ωm and 1250 Ωm. Appropriate depths to which potable water can be obtained from the various locations are recommended in this study.

(Keywords: groundwater, geoelectric, Imo Shale Formation, terrameter, vertical electrical sounding, VES)

## INTRODUCTION

Groundwater is usually referred to as the water beneath the ground surface in auriferous soil pore spaces, or in the fractures of rock formations in form of underground streams. Groundwater is a natural resource with its inherent characteristics being strongly determined by the geologic properties of the host rock.

A porous substratum (rock) that is able to hold and yield (transmit) an appreciable quantity of groundwater upon penetration by a borehole is called an aquifer. Thus, the search for groundwater in an area should be dependent on a reliable empirical knowledge of the local geology of the area.

## <u>Location, Physiography and Geology of the Study Area</u>

The study area lies mainly in the Northern parts of Ehime Mbano Area of Imo State, Southern Nigeria (Figure 1), and falls within the subequatorial belt with average relative humidity values of about 70%; and elevation from about between 61m and 122m above sea level (Egbueri, et al., 2020).

While the wet season spans from Mid-April to the end of October with an annual average rainfall of about 230mm and temperature of the area varies from 29°C to about 33°C.

There are about five different geologic Formations in Imo State of Nigeria but the Imo Shale Formation of the Cenozoic Niger Delta is the study area (Figure 2). The main stratigraphic unit of Imo Shale Formation is Paleocene in age and runs from western to eastern Nigeria in an arcuate belt which is blue-gray in color, and in some places sandy and commonly fossiliferous.

The Formation grades into thick shelly limestone (Ewekoro Formation) in the western part of Nigeria while in the eastern part, it grades in some places into sandy to silty or shaly sandstone (Short and Stauble, 1965). The age of the Formation ranges into Early Eocene and is over-lain by Bende-Ameki Formation.

## Significance of the Study and Choice of Method

Industrialization and urbanization, together with increase in population and rising standards of living usually add pressure on natural resources (Amos-Uhegbu, 2012).

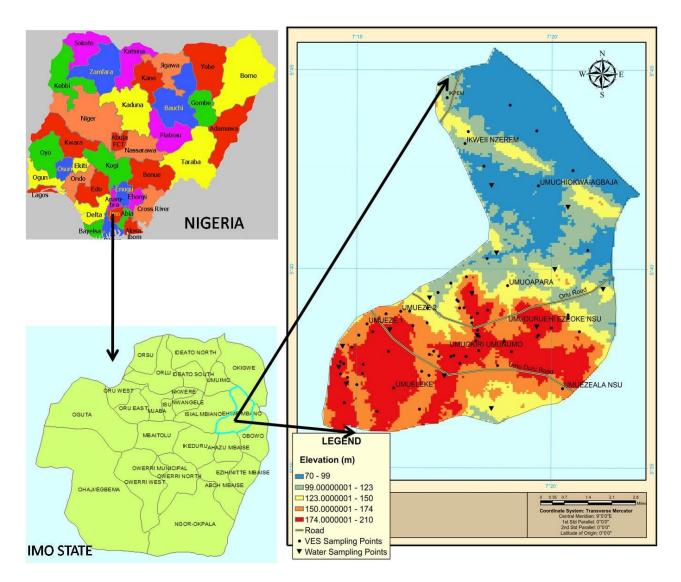


Figure 1: Physiographic Map of the Study area Ehime Mbano Area of Imo State, Nigeria.

The preference of groundwater to surface water; together with the non-availability of the municipal water supply has made groundwater the main source of water supply for almost every sector in the area (Amos-Uhegbu, 2012). The above factors have also added pressure on the groundwater of Ehime-Mbano Area.

Since it has been stated earlier that the search for groundwater in an area should be dependent on a reliable empirical knowledge of the local geology of the area; Imo Shale Formation is well known mostly for its inability to host groundwater, hence the local geology of the area (since it is Imo Shale) need be studied for groundwater potential.

Groundwater studies are usually actualized using modern scientific tools with well outlined techniques.

Geophysical surveys are most widely used because of their basic advantage of providing more accurate results than other methods. Amongst all the geophysical methods, the electrical resistivity method has been the most widely used tool for groundwater investigation. This is because less field manpower is required, and the equipment is portable; hence the field operation is easy. It also has greater depth of penetration thus clarifying the subsurface structure together with the delineation of the groundwater (Amos-Uhegbu, et al, 2017).

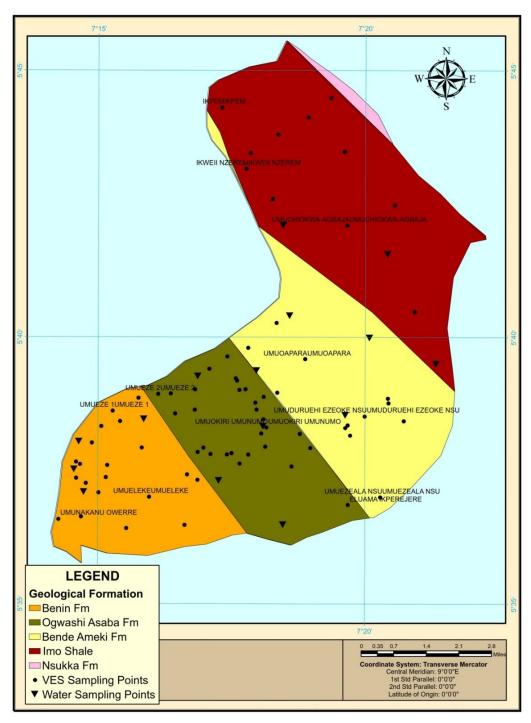


Figure 2: Geologic Map of Ehime Mbano Area of Imo State, Nigeria.

### **MATERIALS AND METHODS**

A total of ten Vertical Electrical Soundings (VES) were acquired using the Schlumberger electrode configuration with two current electrodes 'AB/2' widely spaced out and two potential electrodes 'MN/2'closely spaced in between the current

electrodes all along the survey line. The current electrode spacing 'AB/2' was varied from 1.5 m to a maximum of 350 m; while the potential electrode spacing 'MN/2' was varied from 0.5 m to a maximum of 55 m. Garmin GPS 72 was used in determining the coordinates (longitude, latitude, and elevation height above mean sea

level) of each sounding point; and the ABEM Terrameter SAS 4000 was used in the data acquisition.

A 12V direct current (DC) from a battery was fed into the terrameter which was subsequently passed into the ground through the current electrodes 'AB/2' linked by insulated cables. The resultant potential difference (voltage) was determined using the potential electrodes 'MN/2'.

The observed field data is read off directly from the terrameter, and it is the ratio of the voltage to the current which is a measure of resistance of the subsurface (ground resistance). This measured ground resistance in ohms is used to compute the corresponding apparent resistivity in Ohm-meters by multiplying with the geometric factor.

The value of the geometric factor is a function of electrode spacing, thus giving the required apparent resistivity results as functions of depths of individual layers:

$$\rho a = \pi R((L^2-l^2)/2l)$$
 (1)

Where

ρa = Apparent resistivity.
 L = 'AB/2' = Half current electrode spacing(m).
 I = MN/2 = Half potential electrode spacing (m).

R = Resistance in ohms.

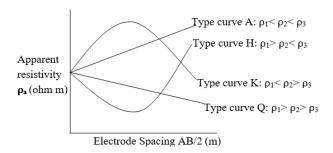
 $\pi R((L^2-l^2)/2l) = Geometric factor (K).$ 

For each sounding point, the subsurface stratigraphy was delineated based on apparent resistivity differences. The apparent resistivity values were plotted against current electrode spacing 'AB/2' on a log-log graph paper to obtain sounding curves. Subsequently, initial estimates of the resistivities and thicknesses of the various geoelectric layers were obtained and used for computer iteration using IPI2Win v. 2.1software package.

#### **RESULTS AND DISCUSSION**

Sounding curve acquired over a horizontally stratified medium is a function of the resistivities and thicknesses of the layers together with the electrode configuration. When the calculated apparent resistivity is plotted against the corresponding half current electrode spacing (AB/2), VES curves are derived, and the letters Q,A,K and H are used singly or in combination to

indicate the variation of resistivity with depth (Figure 3).



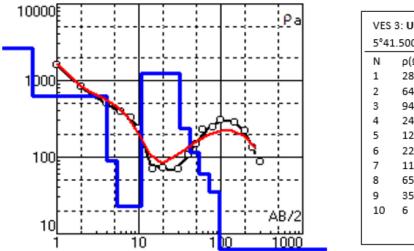
**Figure 3:** Schematic Diagram of Resistivity Type Curves for Layered Structures.

A summary of the interpreted data which is within the limit of the probe is as shown in Table 1. This has revealed the existence of eight to ten geoelectric layers. The topsoil which is the first geoelectric layer has the resistivity varying from  $86~\Omega m$  to  $2850~\Omega m$  with thickness varying from 0.5~m to 0.6~m with 5.5~m only observed at VES 6 station.

Based on the extent (depth) of probe of the survey in the study area, the resistivity type curves, areas of groundwater potential are indicative of single auriferous units. The thickness of the 5<sup>th</sup> layer of VES Stations 1 - 4 ranges from 19m to 21m and they possess favorable groundwater potential from the depth of about 31.5m. Some of the resistivity type curves are as shown (Figures 4, 5, 6, 7, and 8).

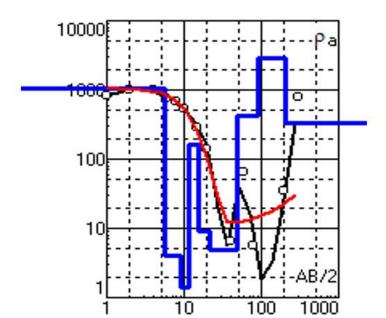
At the vicinity of VES Station 5 after the 6<sup>th</sup> layer at about 57.4 m down to about 187.4 m is a layer indicative of a huge groundwater potential with resistivity of 1300  $\Omega m$ . On the other hand, at VES 6, at about 80 m down to about 187.7 m is also a layer of resistivity of 2900  $\Omega m$  and possesses a favorable huge groundwater potential; while at VES 7 from about 10 m to 28m may likely hold shallow groundwater which may be confined because the shallow nature of the layer may probably rule out the possibility of huge groundwater potential.

The resistivity values of VES Stations 8 -10 do not give favorable groundwater potentials more especially at stations 9 and 10.



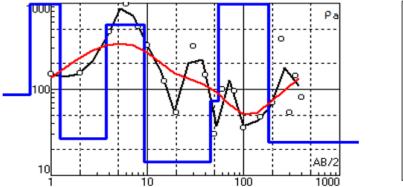
VES 3: Umuosu-Alike Nzerem							
5°41	5°41.500′N, 7°18.493′E, H405ft						
N	ρ(Ωm)	h(m)	d(m)				
1	2800	0.5	0.5				
2	640	3.5	4.0				
3	94	1.5	5.5				
4	24	5.0	10.5				
5	1250	20.0	30.5				
6	225	10.5	41.0				
7	118	10.5	51.5				
8	65	18.0	69.5				
9	35.5	21.0	90.5				
10	6						

Figure 4: Typical Vertical Electrical Sounding Curve of VES 3.



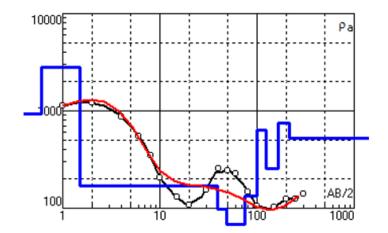
VES	VES 6: Umuemeke Agbaja								
5°4	5°41.706'N, 7°17.085'E, H347ft								
N	N ρ(Ωm) h(m) d(m)								
1	1150	5.5	5.5						
2	4.0	3.5	9.0						
3	1.4	3.0	12.0						
4	180	4.0	16.0						
5	9.2	6.0	22.0						
6	4.2	25.0	47.0						
7	430	35.0	82.0						
8	2900	130.0	212.0						
9	320								

**Figure 5:** Typical Vertical Electrical Sounding Curve of VES 6.



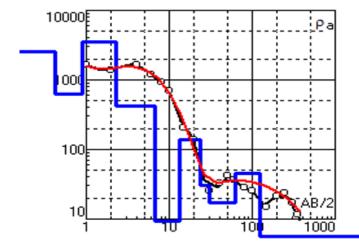
VES 5: <b>Umuchukwu Mgbam</b> <b>Agbaja</b> 5°41.679'N, 7°14.812'E, H350ft								
1	86	0.6	0.6					
2	1150	0.6	1.2					
3	27	2.6	3.8					
4	580	5.6	9.4					
5	15	38.0	47.4					
6	75	10.0	57.4					
7	1300	130.0	187.4					
8	24							

Figure 6: Typical Vertical Electrical Sounding Curve of VES 5.



	VES 8: <b>Agbaja</b> 5°40.756'N, 7°15.482'E, H502ft							
N	<del></del>							
1	930	0.6	0.6					
2	2850	1.0	1.6					
3	170	38.4	40.0					
4	95	4.8	44.8					
5	62	9.0	53.8					
6	140	25.0	78.8					
7	650	26.0	104.8					
8	245	45.0	149.8					
9	720	50.0	199.8					
10	540							

Figure 7: Typical Vertical Electrical Eounding Curve of VES 8.



	VES 9: Umudike Umunuhu-Nsu 5°39.136'N, 7°20.635'E, H205ft					
· · · · · · · · · · · · · · · · · · ·						
N 1	ρ(Ωm) 2650	h(m) 0.5	d(m) 0.5			
2	640	0.4	0.9			
3	3500	1.6	2.5			
4	430	4.0	6.5			
5	9.5	7.0	13.5			
6	140	9.0	22.5			
7	31	6.0	28.5			
8	18	38.0	66.5			
9	45	71.0	137.5			
10	4					

Figure 8: Typical Vertical Electrical Sounding Curve of VES 9.

Table 1: A Summary of the Interpreted VES Data and their Locations.

VES	Location	GPS	Resistivity	Inferred	Thickness	Maximum	Layer	Туре
Station		reading	of	Lithology of	of layers	depth of	conductivity σ	Curves
			layers (Ωm)	layers	(m)	layers (m)		
1	Umuna –	5°40.914'N,	ρ1 =2600	SandyTopsoil	t1 =0.5	h1 =0.5	σ1 =0.000385	QQKQQQQQ
	Nzerem	7°19.954'E,	ρ2 =650	Silt	t2 = 3.5	h2 = 4.0	σ2 =0.001538	
		H405ft	ρ3 =100	Clay	t3 = 1.5	h3 = 5.5	σ3 =0.010000	
			ρ4 = 25	Clay	t4 = 5.0	h4 = 10.5	σ4 =0.008000	
			ρ5 =1250	Sand	$t_5 = 21.0$ $t_6 = 8.5$	h5 = 31.5 h6 = 40.0	σ5 =0.000800 σ6 =0.004167	
			ρ6 =240 ρ7 =110	Clay	t7 = 15.0	h7 = 55.0	σ7 =0.009090	
			ρ8 =60	Clay	t8 =20.5.	h8 = 76.0	σ8 =0.016667	
			ρθ = 00	Clay Clay	t9 =24.0.	h9 =100.0	σ9 =0.025000	
			ρ10= 10	Clay	t10 = ?	h10 = ?	σ10=0.100000	
•	Nzerem	5°40.438'N,	ρ1 =2400	SandyTopsoil	t1 =0.5	h1 =0.5	σ1 =0.000417	QQHKQQQQ
2	Market Road	7°20.755'E,	$\rho$ 2 =600	Silt	$t_1 = 0.5$ $t_2 = 3.5$	h2 = 4.0	$\sigma^2 = 0.001667$	QQIINQQQQ
		H404ft	ρ3 =95	Clay	t3 = 1.5	h3 = 5.5	σ3 =0.010526	
			ρ4 = 24	Clay	t4 = 5.5	h4 = 11.0	σ4 =0.041667	
			ρ5 =1240	Sand	t <sub>5</sub> = 19.0	h5 = 30.0	σ5 =0.000806	
			ρ6 =240	Clay	$t_6 = 8.5$	h6 = 40.5	σ6 =0.004167	
			ρ7 =110	Clay	t7 = 15.0	h7 = 55.5	σ7 =0.009090	
			ρ8 =60	Clay	t8 =20.5.	h8 = 76.0	$\sigma 8 = 0.016667$	
			$\rho 9 = 40$	Clay	t9 =24.0.	h9 =100.0	$\sigma$ 9 =0.025000	
			ρ10= 9	Clay	t10 = ?	h10 = ?	σ10=0.111111	
3	Umuosu-	5°41.500'N,	ρ1 =2800	SandyTopsoil	t1 =0.5	h1 =0.5	σ1 =0.000357	QQHKQQQQ
3	Alike Nzerem	7°18.493'E,	ρ2 =640	Silt	t2 = 3.5	h2 = 4.0	$\sigma$ 2 =0.001563	
		H405ft	ρ3 =94	Clay	t3 = 1.5	h3 = 5.5	$\sigma$ 3 =0.010638	
			$\rho 4 = 24$	Clay	t4 = 5.0	h4 = 10.5	$\sigma$ 4 =0.041667	
			ρ5 =1250	Sand	$t_5 = 20.0$	h5 = 30.5	σ5 =0.000800	
			ρ6 =225	Clay	t <sub>6</sub> = 10.5	h6 = 41.0	σ6 =0.004444	
			ρ7 =118	Clay	t7 = 10.5	h7 = 51.5	σ7 =0.008475	
			ρ8 =65	Clay	t8 = 8.0.	h8 = 69.5	σ8 =0.015385	
			$\rho 9 = 35.5$	Clay	t9 =21.0. t10 = ?	h9 =90.5 h10 = ?	σ9 =0.028169 σ10=0.166667	
			ρ10= 6	Clay				
4	Alike Nzerem	5°40.111'N,	ρ1 =2650	SandyTopsoil	t1 =0.5	h1 = 0.5	σ1 =0.000377	QQHKQQQQ
		7°22.051'E, H406ft	ρ2 =650	Silt	t2 = 3.4	h2 = 3.9	σ2 =0.001538	
		П400П	ρ3 =95	Clay	t3 = 1.4 t4 = 4.8	h3 = 5.3	σ3 =0.010526	
			ρ4 = 23	Clay		h4 = 10.1 h5 = 31.1	σ4 =0.043478 σ5 =0.000800	
			ρ5 =1250	Sand	$t_5 = 21.0$ $t_6 = 11.5$	h6 = 42.6	σ6 =0.004000	
			ρ6 =250 ρ7 =120	Clay	t7 = 15.0	h7 = 51.5	σ7 =0.008333	
			ρ8 =63	Clay	t8 =25.0.	h8 = 76.5	σ8 =0.015873	
			$\rho 9 = 36$	Clay	t9 =40.0.	h9 =116.5	σ9 =0.027778	
			ρ10= 7	Clay	t10 = ?	h10 = ?	σ10=0.142857	
	Umuchukwu	5°41.679'N,	ρ1 = 36	Clay ClayeyTopsoil	t1 =0.6	h1 =0.6	σ1 =0.027779	KHKHQK
5	Mgbam	7°14.812'E,	$\rho 1 = 30$ $\rho 2 = 1150$	Sand	t2 = 0.6	h2 = 1.2	$\sigma = 0.027779$ $\sigma = 0.000870$	MINION
	Agbaja	H350ft	ρ3 =27	Clay	t3 = 2.6	h3 = 3.8	σ3 =0.037037	
	• •		$\rho = 580$	Silt	t4 = 5.6	h4 = 9.4	σ4 =0.001724	
	1		ρ5 =15	Clay	t <sub>5</sub> = 38.0	h5 = 47.4	σ5 =0.066667	
	1		ρ6 =75	Clay	t <sub>6</sub> = 10.0	h6 = 57.4	σ6 =0.013333	
			ρ7 =1300	Sand	t7=130.0	h7 = 187.4	σ7 =0.000769	
			ρ8 =24	Clay	t8 = ?	h8 = ?	σ8 =0.041667	
6	Umuemeke	5°41.706'N,	ρ1 =1150.0	SandyTopsoil	t1 =5.5	h1 = 5.5	σ1 =0.000870	QHKQHAK
0	Agbaja	7°17.085'E,	ρ2 =4.0	Clay	t2 = 3.5	h2 = 9.0	σ2 =0.250000	
		H347ft	ρ3 =1.4	Clay	t3 = 3.0	h3 = 12.0	σ3 =0.714286	
			ρ4 = 180.0	Clay	t4 = 4.0	h4 = 16.0	σ4 =0.005556	
	1		ρ5 =9.2	Clay	$t_5 = 6.0$	h5 = 20.0	σ5 =0.108696	
	1		ρ6 =4.2	Clay	$t_6 = 25.0$	h6 = 45.0	σ6 =0.238095	
			ρ7 =430.0	Silt	t7 = 35.0	h7 = 80.0	σ7 =0.002326	
	1		ρ8 =2900.0	Sand	t8=130.0	h8 = 210.0	σ8 =0.000345	
			$\rho$ 9 = 320	Clay	t9 = ?	h9 = ?	σ9 =0.003125	

Volume 23. Number 1. May 2022 (Spring)

7	Umuagim	5°41.275'N,	ρ1 =2800	SandyTopsoil	t1 = 0.5	h1 = 0.5	σ1 =0.000357	QQHKQQQQ
7	Nzerem	7°20.559'E,	ρ1 =2600	Silt	t1 = 0.5 t2 = 3.4	h2 = 3.9	$\sigma^2 = 0.000357$	QQHNQQQQ
	NZCICIII	H405ft	ρ2 =040	Clay	t3 = 1.2	h3 = 5.1	σ3 =0.010638	
			$\rho 3 = 94$ $\rho 4 = 24$	Clay	t3 = 1.2 t4 = 4.8	h4 = 9.9	σ4 =0.041667	
			ρ5 =1300	Sand	$t_5 = 19.0$	h5 = 28.9	σ5 =0.000769	
			ρ6 =240	Clay	$t_6 = 10.0$	h6 = 38.9	σ6 =0.004167	
			ρ7 =130	Clay	t7 = 13.0	h7 = 51.9	σ7 =0.007692	
			ρ8 =58		t8 =19.0.	h8 = 70.9	σ8 =0.017241	
			$\rho 9 = 37$	Clay	t9 = 4.5.	h9 = 75.4	σ9 =0.027027	
			$\rho = 37$ $\rho = 5$	Clay	t10 = ?	h10 = ?	σ10=0.200000	
	A b i -	F940 7FCINI		Clay SiltyTopsoil				KOOLIAKUK
8	Agbaja	5°40.756'N, 7°15.482'E,	ρ1 =930	Sand	t1 = 0.6 t2 = 1.0	h1 = 0.6 h2 = 1.6	σ1 =0.001075 σ2 =0.000351	KQQHAKHK
		H502ft	ρ2 =2850		t2 = 1.0 t3 = 38.4	h3 = 40.0	$\sigma_3 = 0.000351$	
		1130210	ρ3 =170	Clay	t4 = 4.8	h4 =44.8	$\sigma 4 = 0.010526$	
			ρ4 = 95	Clay	$t_5 = 9.0$	h5 = 53.8	$\sigma 5 = 0.016129$	
			ρ5 =62	Clay	$t_6 = 25.0$	h6 = 78.8	$\sigma 6 = 0.007143$	
			ρ6 =140	Clay	$t_0 = 25.0$ $t_7 = 26.0$	h7 = 104.8	σ7 =0.001538	
			ρ7 =650	Silt	t8 =45.0.	h8 = 149.8	σ8 =0.004082	
			ρ8 =245	Clay	t9 = 50.	h9 = 199.8	σ9 =0.001389	
			$\rho 9 = 720$	Silt	t10 = ?	h10 = ?	σ10=0.001852	
			ρ10= 540	Silt				
9	Umudike Umunuhu-	5°39.136'N,	ρ1 =2650	SandyTopsoil	t1 = 0.5	h1 = 0.5	σ1 =0.000377	HKAHKQHK
	Nsu	7°20.635'E, H205ft	ρ2 =640	Silt	t2 = 0.4	h2 = 0.9	σ2 =0.001563	
	NSU	пиозії	ρ3 =3500	Sand	t3 = 1.6	h3 = 2.5	σ3 =0.000286	
			$\rho 4 = 430$	Silt	t4 = 4.0	h4 =6.5	σ4 =0.002326 σ5 =0.105263	
			ρ5 =9.5	Clay	$t_5 = 7.0$ $t_6 = 9.0$	h5 = 13.5	$\sigma 6 = 0.007143$	
			ρ6 =140	Clay	$t_6 - 9.0$ t7 = 6.0	h6 = 22.5 h7 = 28.5	σ7 =0.032258	
			ρ7 =31	Clay	t8 = 38.0.	h8 = 66.5	σ8 =0.055556	
			ρ8 =18	Clay	t9 =71.0.	h9 = 137.5	σ9 =0.022222	
			$\rho 9 = 45$	Clay	t10 = ?	h10 = ?	σ10=0.250000	
	ļ.,		ρ10= 4	Clay				
10	Umuezeala-	5°37.026'N,	ρ1 =590	SiltyTopsoil	t1 = 0.6	h1 = 0.6	σ1 =0.001695	KQHAAAAA
	Uhu Road	7°15.571'E, H692ft	ρ2 =1250	Sand	t2 = 4.0	h2 = 4.6	σ2 =0.000800	
		H092π	$\rho 3 = 27.5$	Clay	t3 = 3.0	h3 = 7.6	σ3 =0.036364	
			$\rho 4 = 7.2$	Clay	t4 = 14.6	h4 = 22.2	σ4 =0.138889	
			ρ5 =28.5	Clay	t <sub>5</sub> = 70.4	h5 = 92.6	σ5 =0.035088	
			ρ6 =29.5	Clay	$t_6 = 29.4$	h6 = 122.0	σ6 =0.033898	
			ρ7 =30.5	Clay	t7 = 40.6 t8 =43.0.	h7 = 162.6	σ7 =0.024631	
			ρ8 =32.5	Clay	t8 =43.0. t9 =47.0.	h8 = 205.6 h9 = 252.6	σ8 =0.030769 σ9 =0.022222	
			$\rho 9 = 45$	Clay	t9 =47.0. t10 = ?	n9 = 252.6 h10 = ?	σ10=0.022222 σ10=0.014599	
			ρ10= 68.5	Clay	110 - !	1110 - !	010-0.014099	

### **CONCLUSION**

The investigation of the groundwater potential of the Imo shale Formation in Ehime-Mbano Area using the electrical resistivity technique in this study has shown that groundwater in the form of confined aquifer is likely obtainable between an average depth of 31 m and 40 m within the vicinity of VES 1-4 (Nzerem 1- 4). The resistivity of the aquifer layer ranged between 1240  $\Omega m$  and 1250  $\Omega m$ . The study has also shown that groundwater can be sourced from the seventh layer in VES 5 and the eight layer in VES 6 at Agbaja 1 and 2, respectively.

Based on the depth of investigation, the possibility of obtaining sustainable groundwater at the vicinity of VES 9 and 10 is not recommended.

Finally, the study has indicated that the entire area may not have huge sustainable groundwater potential. This is because; most of the potential groundwater host rocks are single-layered not multi-layered.

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